

**A comparison of evapotranspiration rates for willow and
reed in a riverine fen**

First measurement results

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financed by:



Table of contents

- Introduction

- description of the investigation area**
- occasion of the investigation**

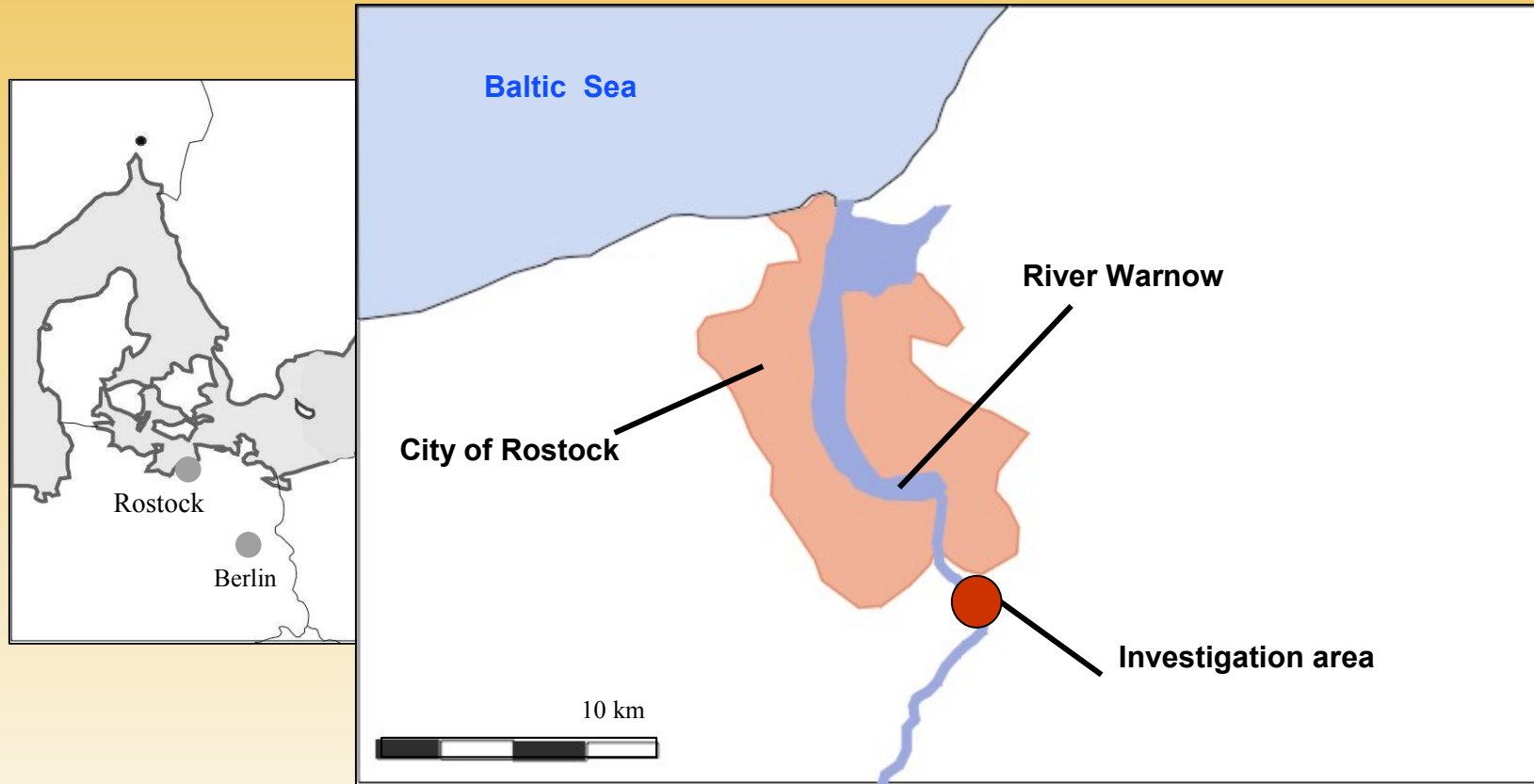
- Method

- meteorological measurement program**
- hydrological measurement program**

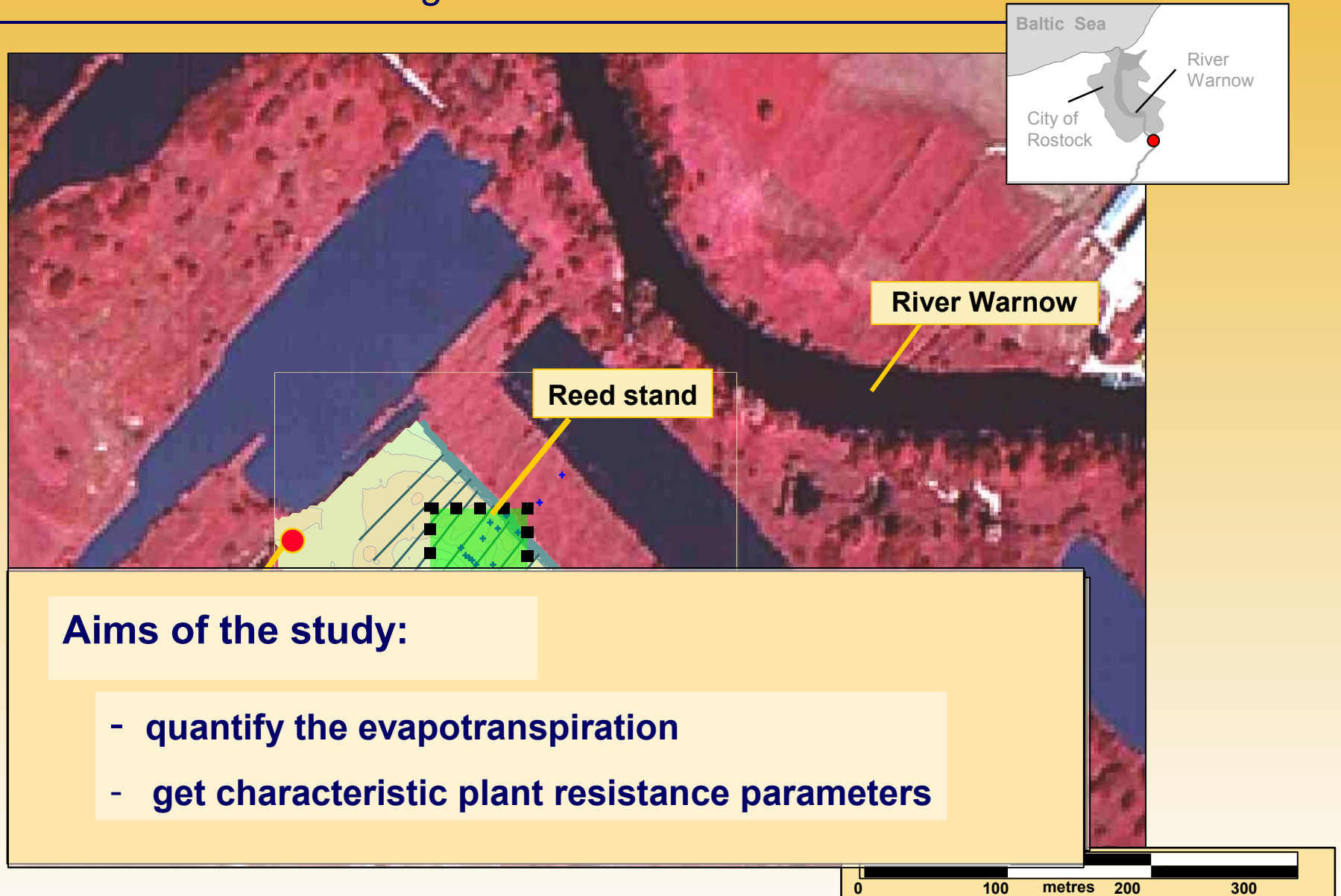
- Results

Introduction – The investigation area

Situation of the area



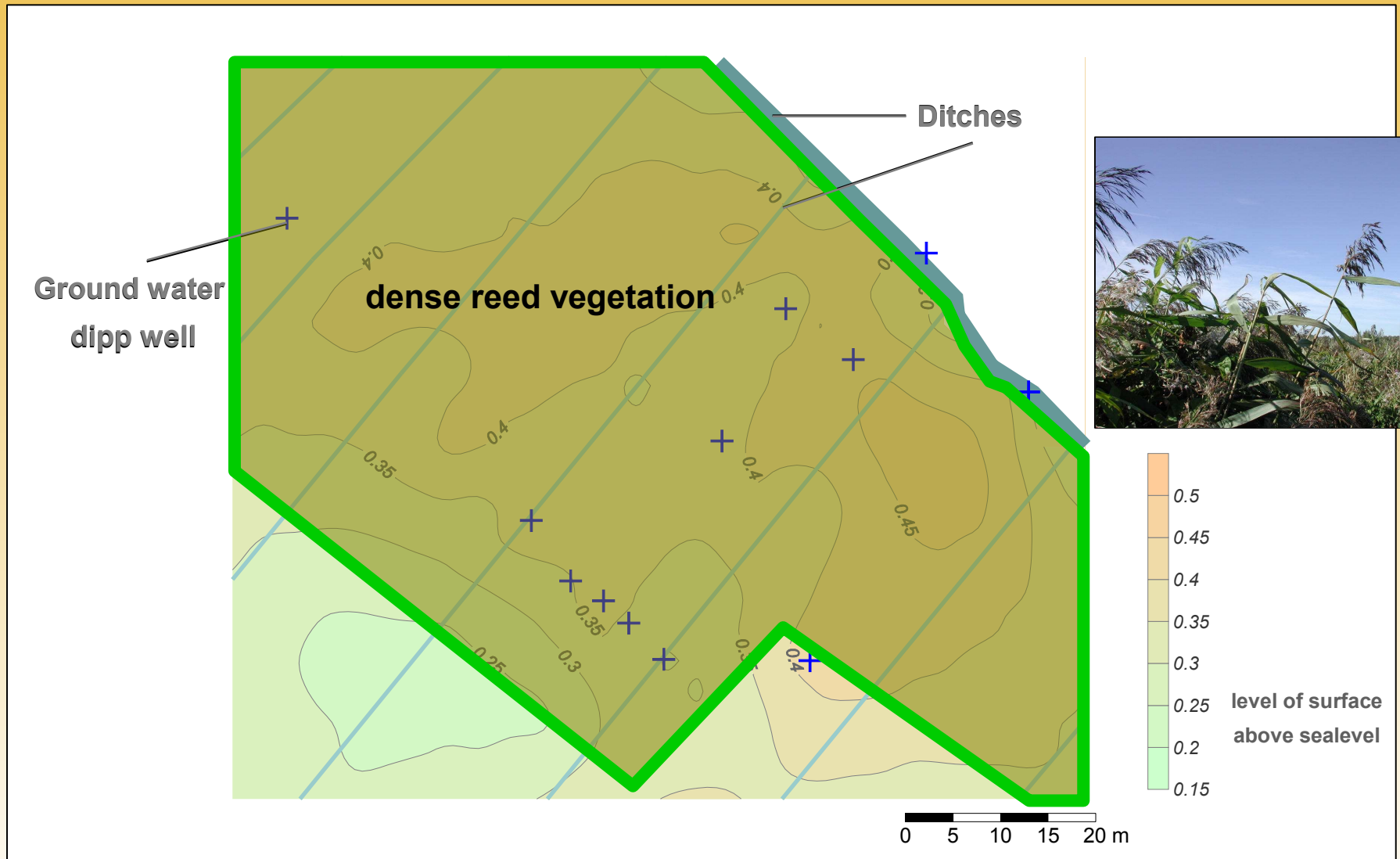
Introduction – The investigation area



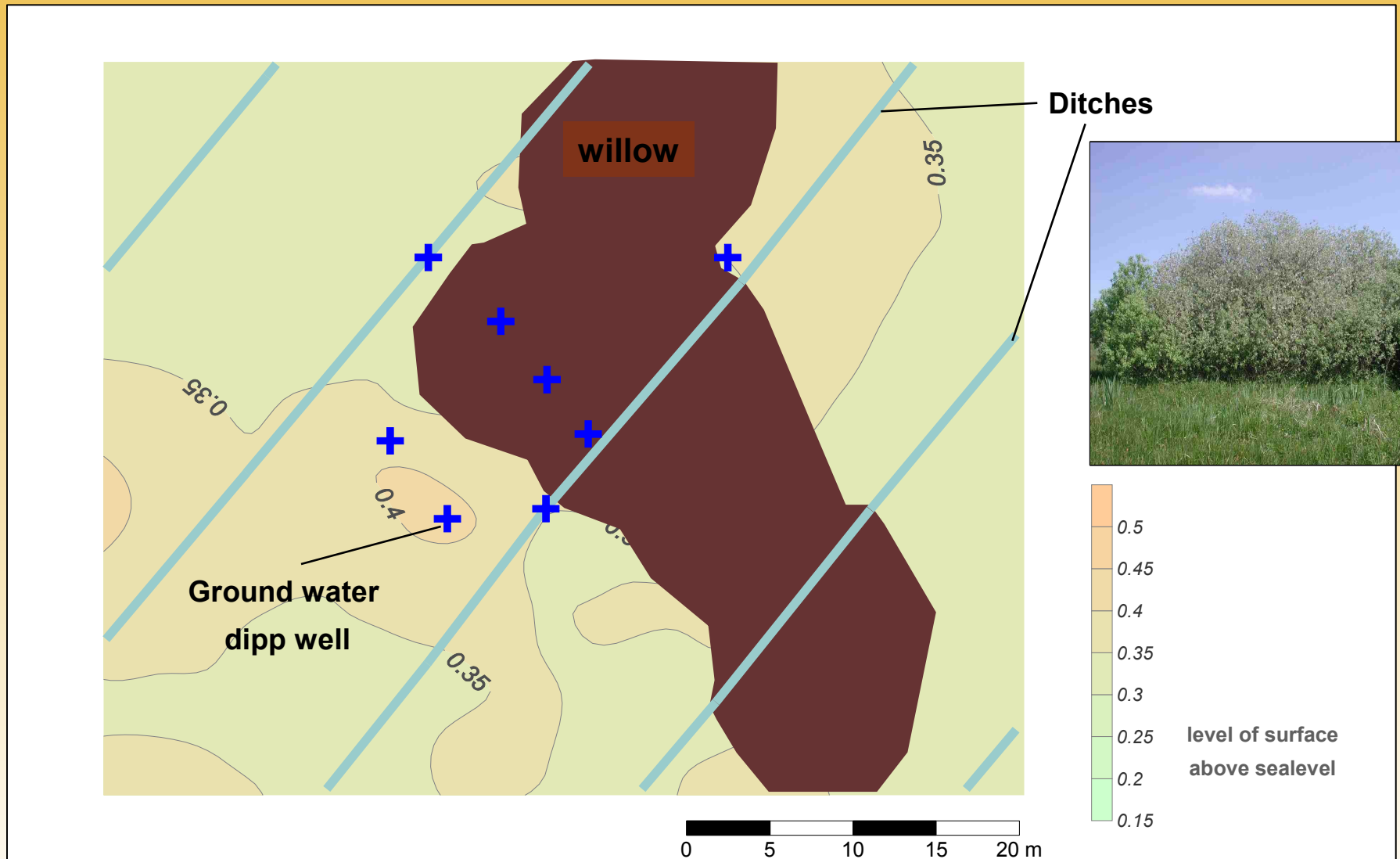
Aims of the study:

- quantify the evapotranspiration
- get characteristic plant resistance parameters

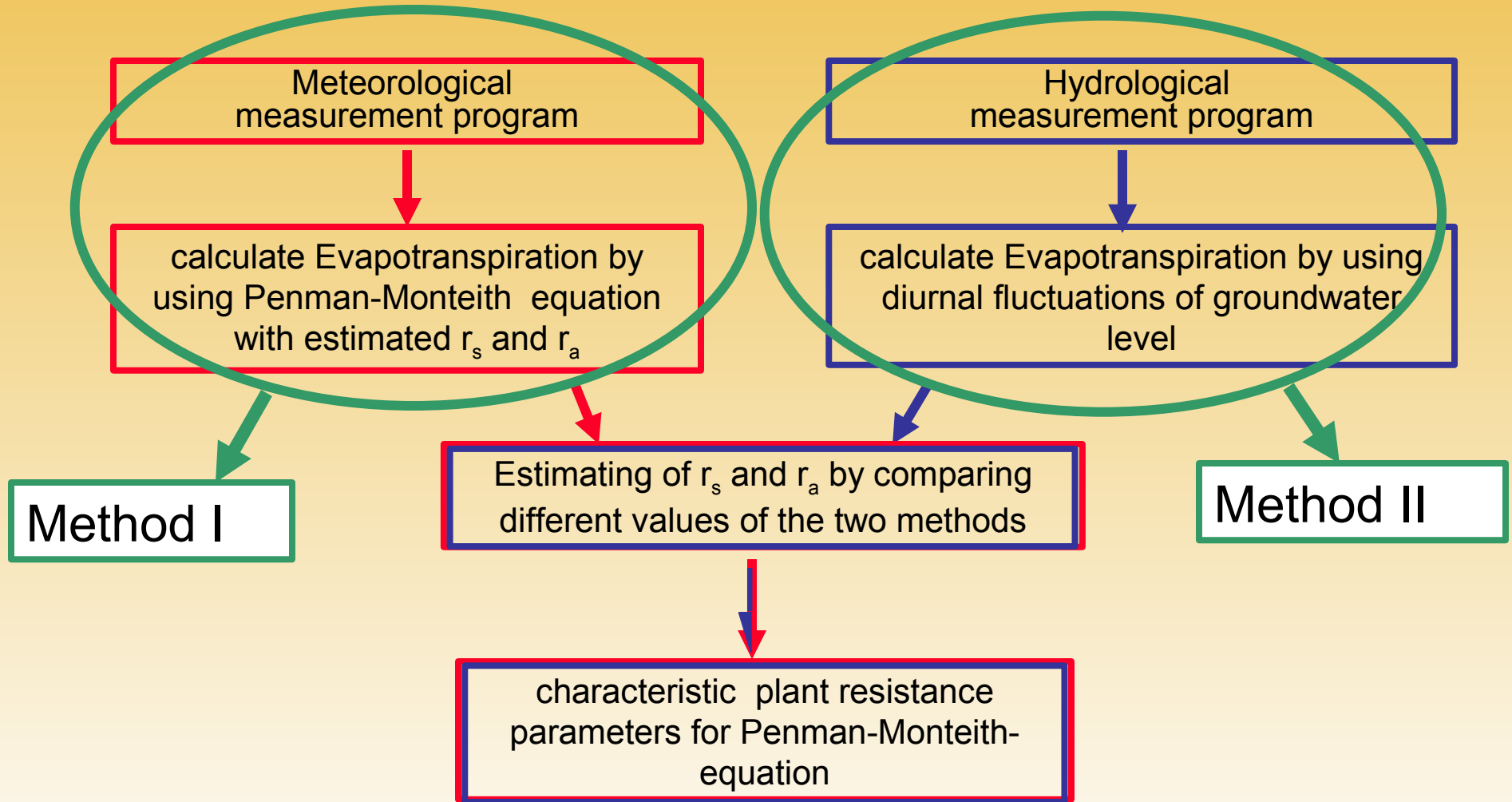
The reed stand



The willow stand



Method



Method I – Meteorological measurement program

Penman - Monteith - equation

$$ETa = \frac{1}{L^*} \times \frac{s \times (Rn - G) + \frac{\rho \times c_p}{r_a} \times (e_s(T) - e)}{s + \gamma \times \left(1 + \frac{r_s}{r_a}\right)}$$

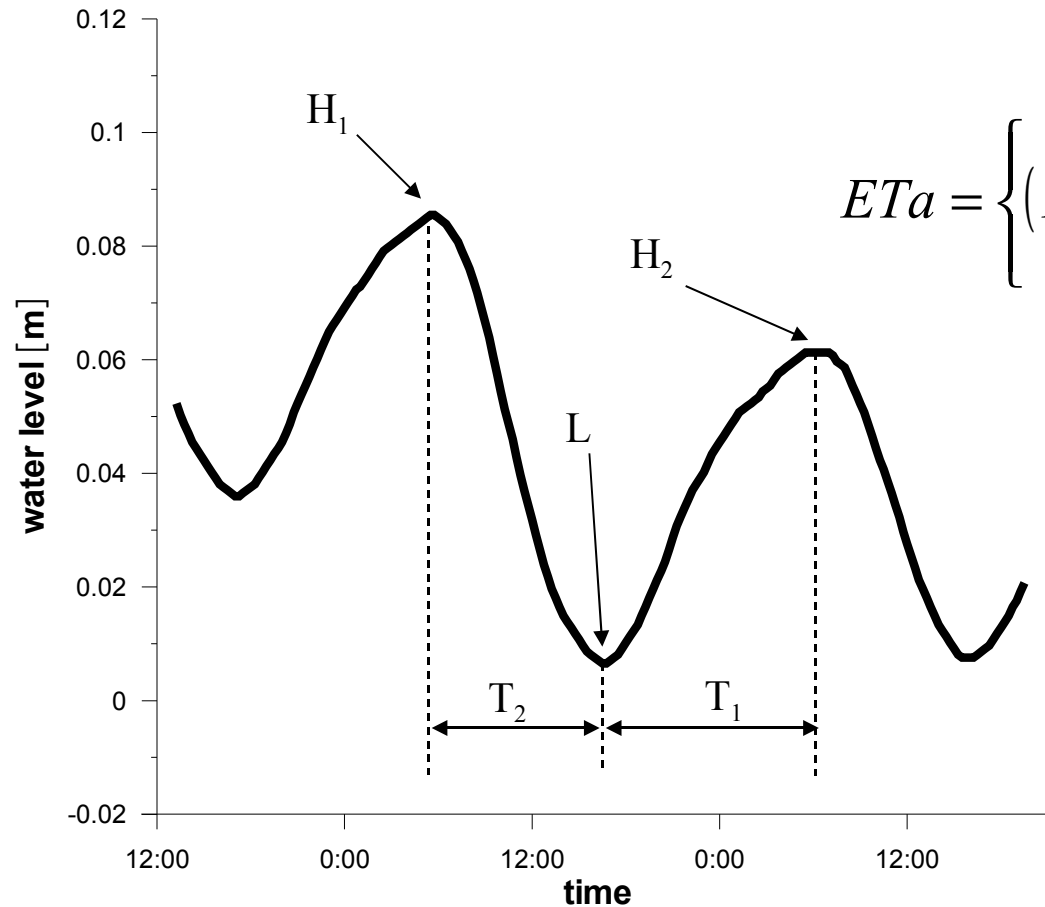
- **constant factors** : air density ρ , specific heat of the air c_p ,
psychrometric constant γ

- **meteorological variables** → measured

- **botanic/meteorological variables** → estimated parameters
for reed and willow

Method II – Hydrological measurement program

Using of *diurnal ground water level fluctuation* to calculate the daily evapotranspiration rate



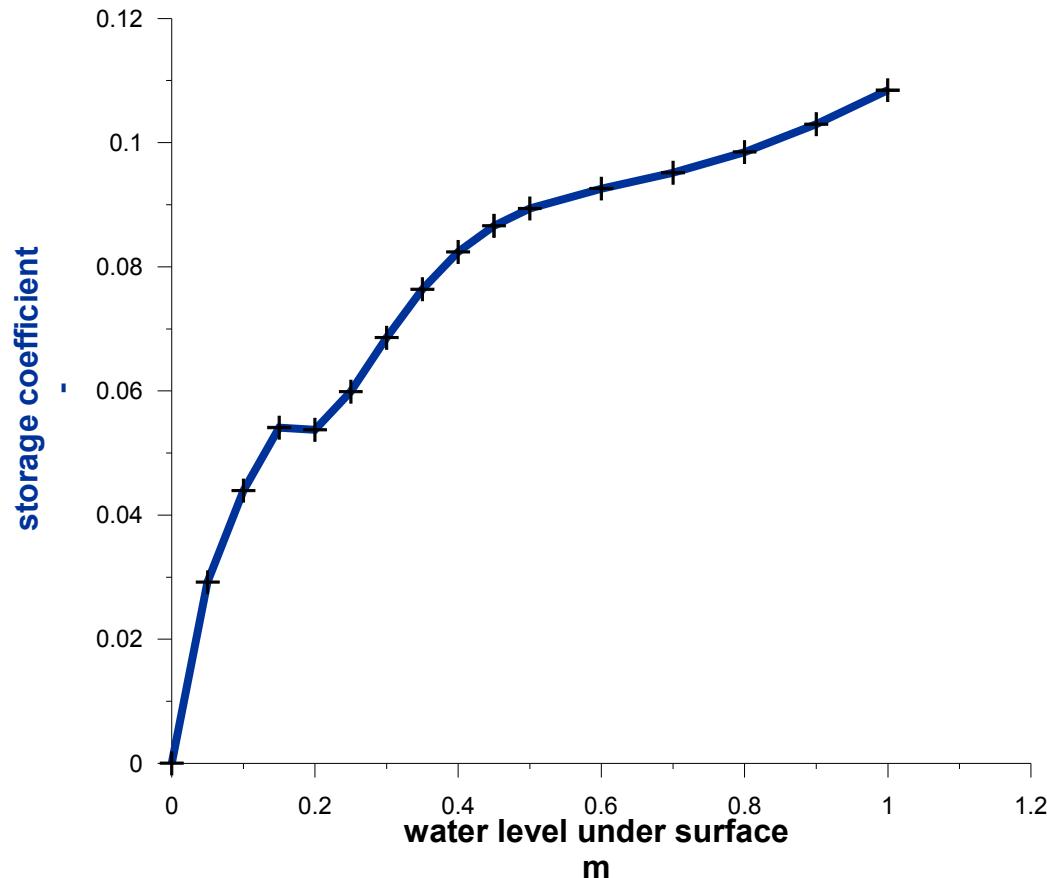
$$ETa = \left\{ (H_1 - L) + \left[(H_2 - L) \cdot \frac{T_1}{T_2} \right] \right\} \cdot S$$

„Draw down recharge method“
(Hays, 2003)

S = storage coefficient

Storage coefficient function

The storage coefficient – a *non constant* value



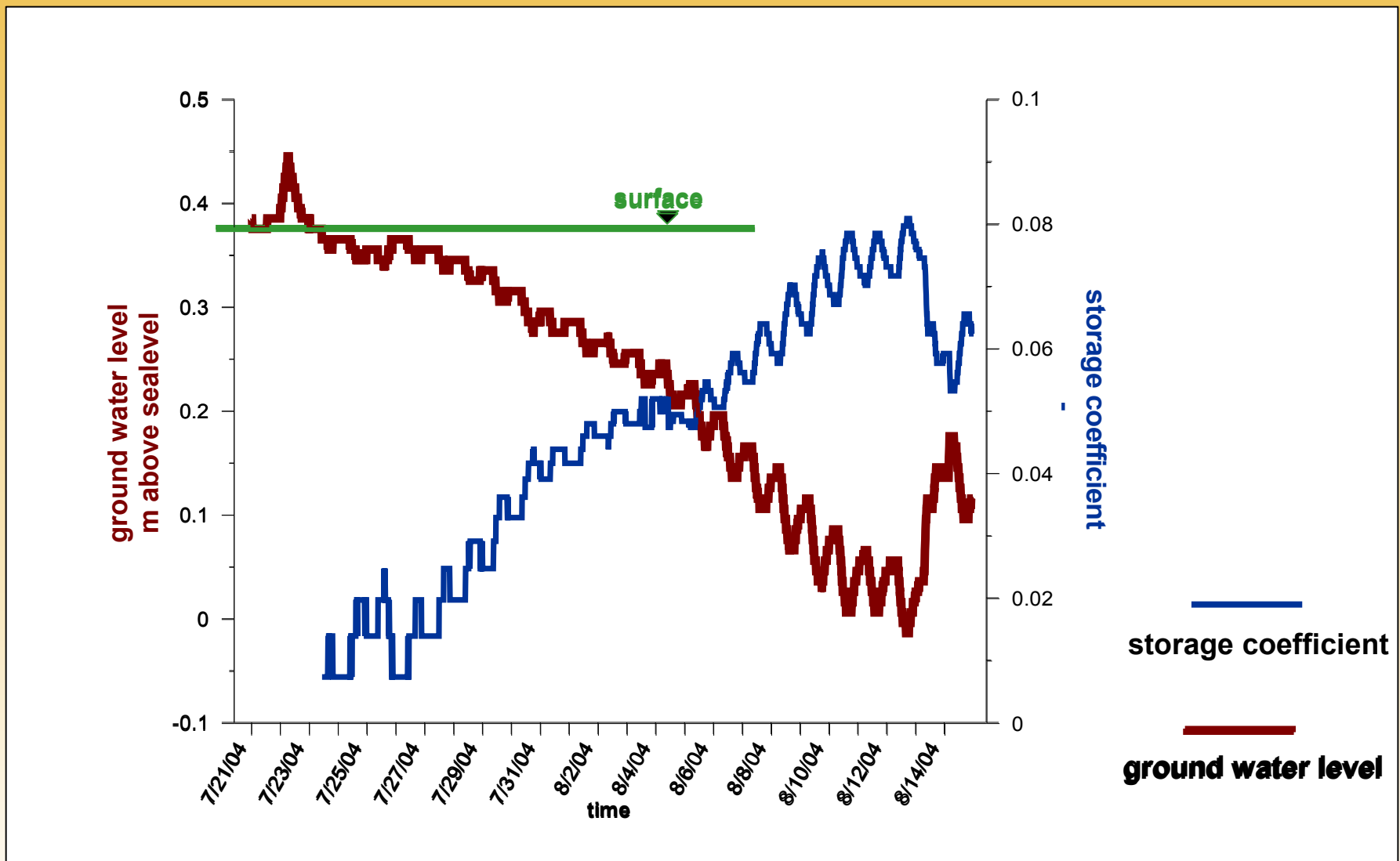
water level under surface (m)	storage coefficient (-)
0.00	0.000
0.05	0.029
0.10	0.044
0.15	0.054
0.20	0.054
0.25	0.060
0.30	0.069
0.35	0.076
0.40	0.082
0.45	0.087
0.50	0.089
0.60	0.093
0.70	0.095
0.80	0.098
0.90	0.103
1.00	0.108

(Trübger 2005)

Including:

- **dual porosity** (based on Durner, 1994)
- **Van Genuchten parameters**

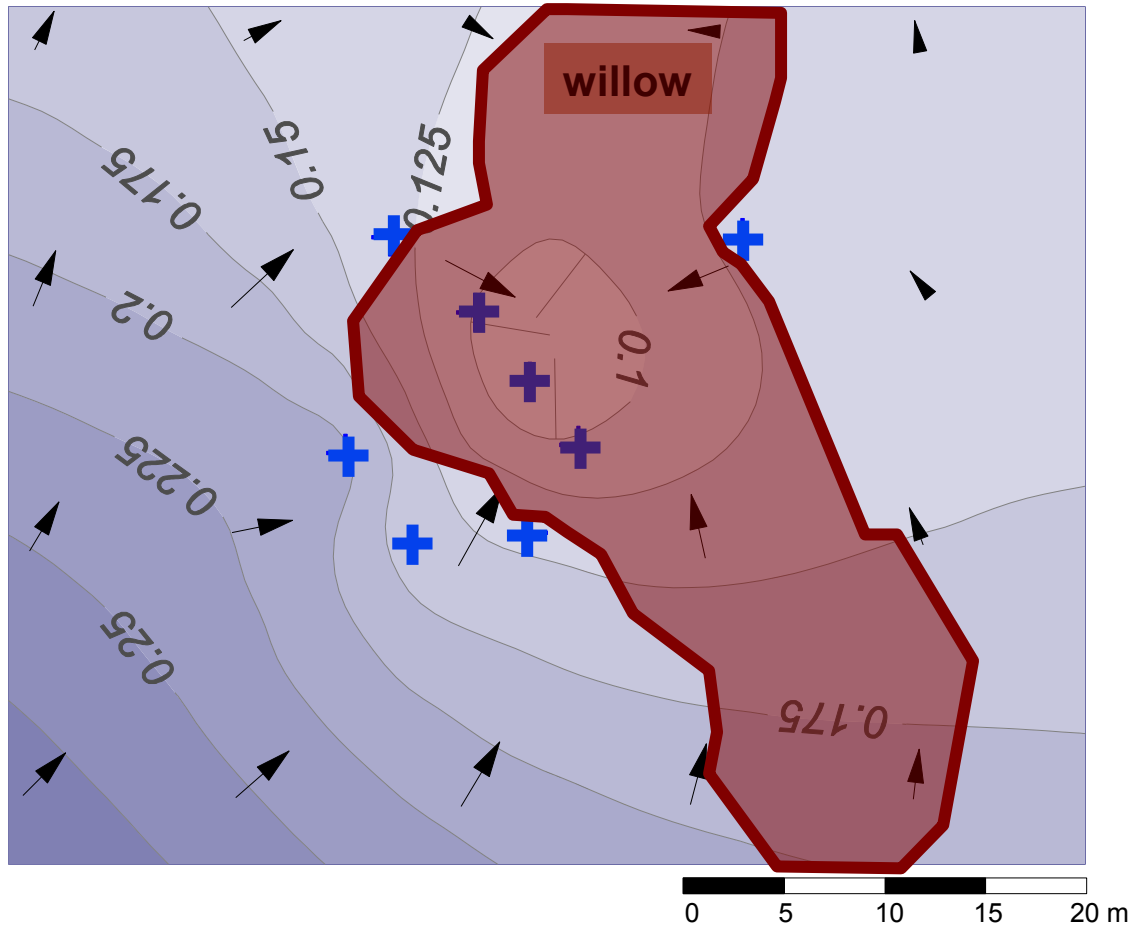
An application of the storage coefficient function



Results – The hydrologic situation on 09.08.2004

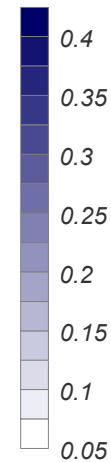
- at the willow stand

→ ~~no significant depression of groundwater~~
to the central willow



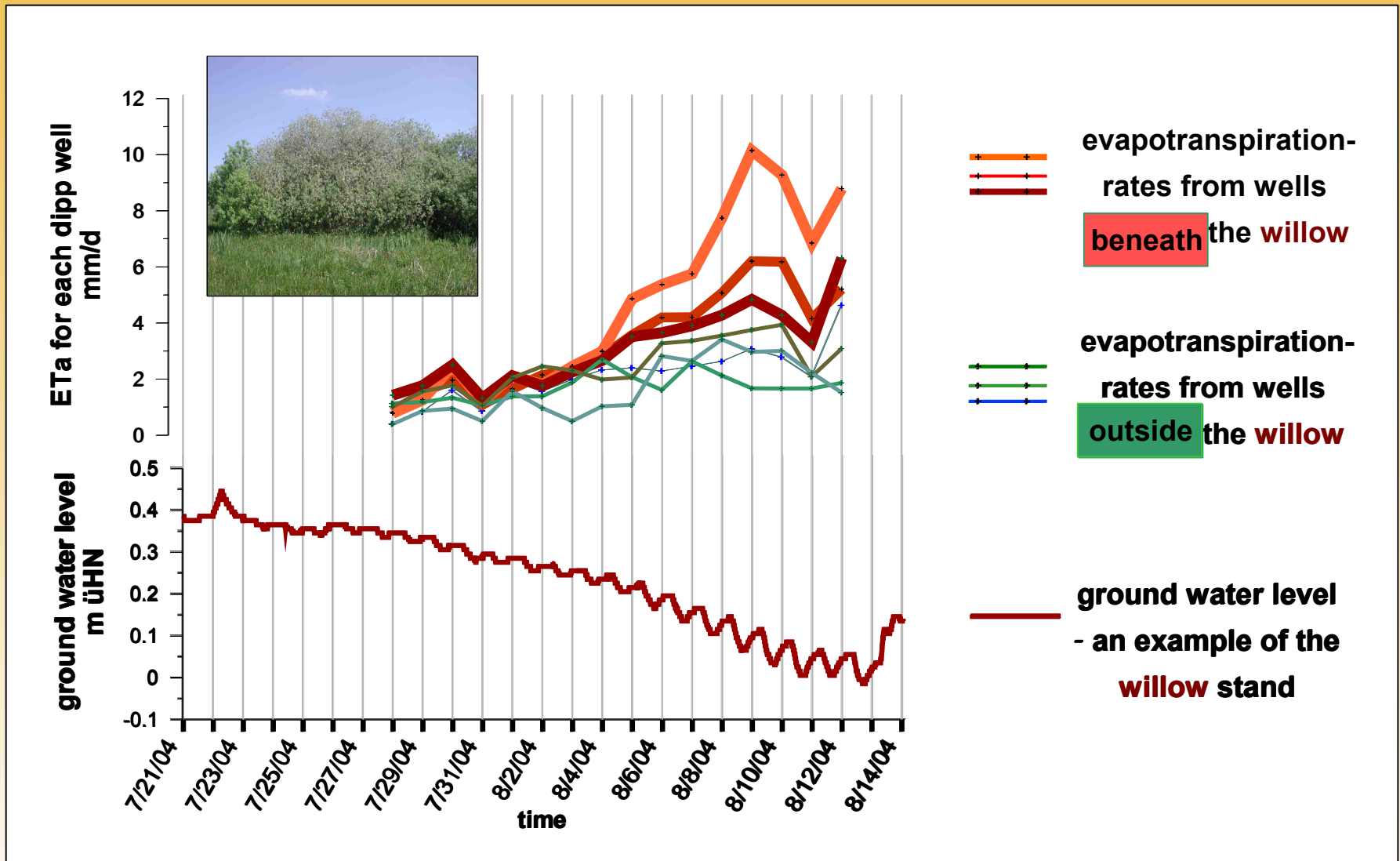
+ ground water dipp well

↗ flow direction



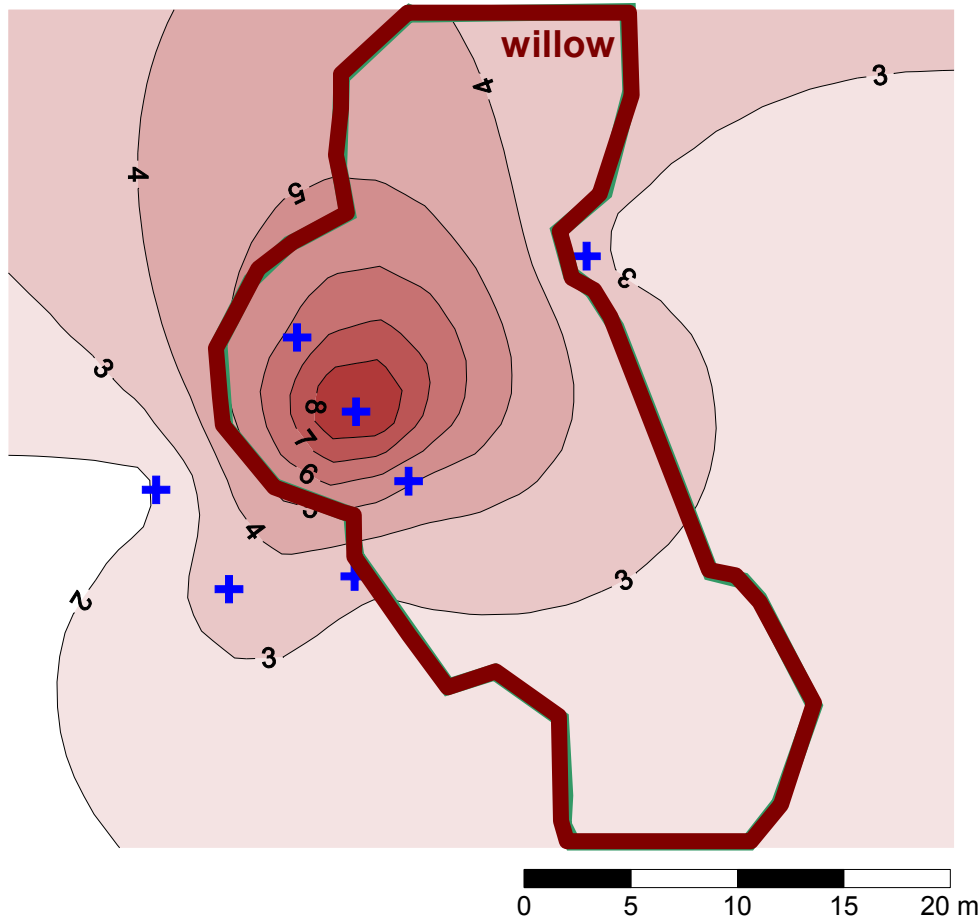
ground water level
m above sealevel

Results – The Draw down recharge method – at the *willow* stand

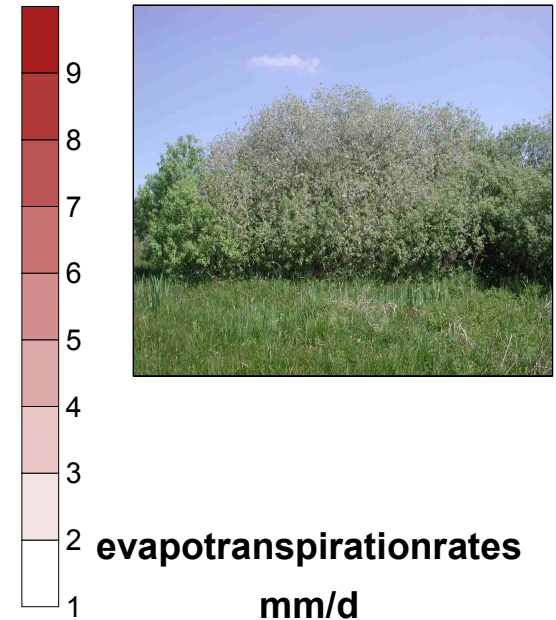


Results – The Draw down recharge method – at the *willow* stand

➔ maximum Eta in the central willow

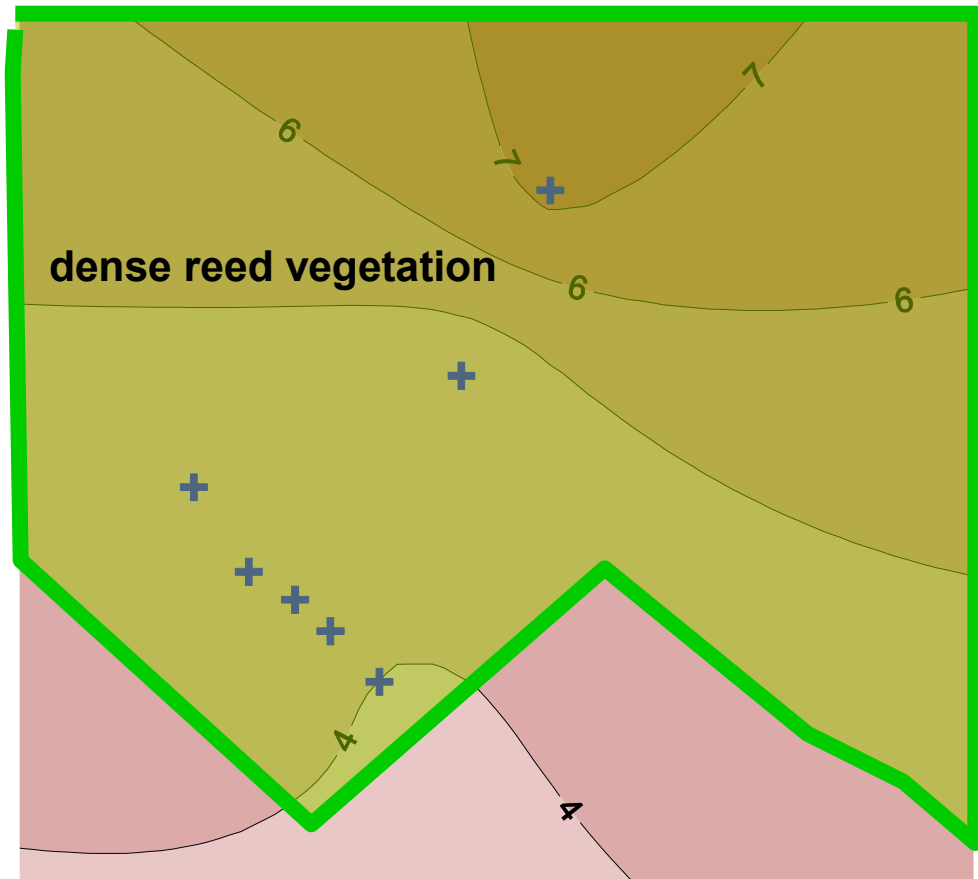


⊕ ground water dipp well

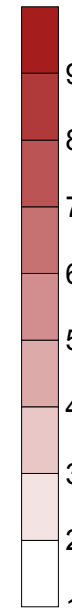


Results – The Draw down recharge method – at the *reed* stand

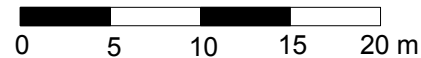
→ uniform distribution of Eta at the reed stand



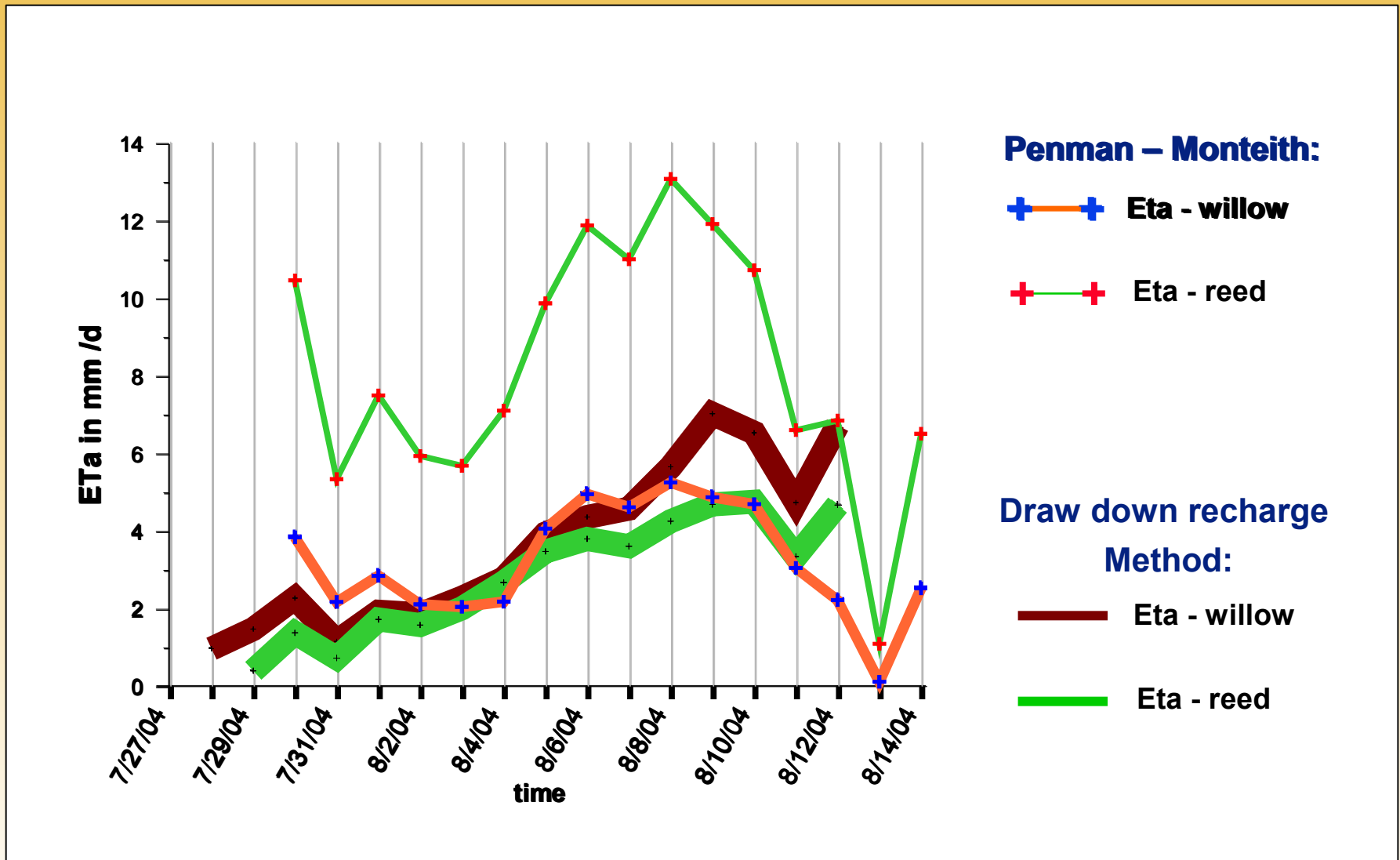
+ ground water dip well



evapotranspiration rates
mm/d



Results – A comparison



Final remarks

- this estimation shows differences between the Eta rates for willow and reed in the following areas (based on diurnal ground water fluctuations):

- cone of depression of ground water beneath the willow

- the area distribution of Eta at the two stands

- max Eta_{willow} = 10,0 mm/d - max Eta_{reed} = 7,0 mm/d

- cf. previous literature: 8,8 to 10,0 mm/d*

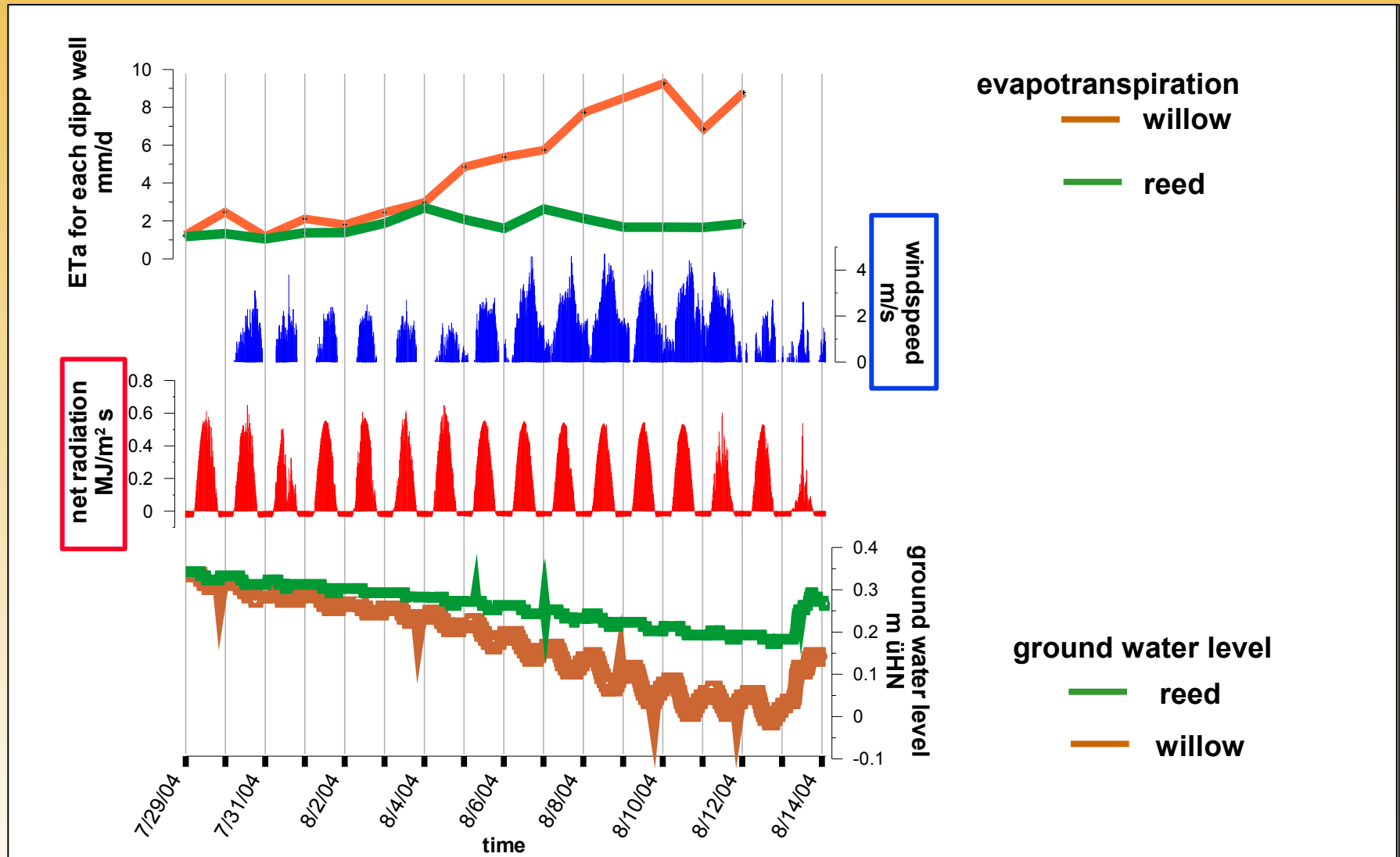
- 7,0 to 10,0 mm/d*

- the using of Penman-Monteith equation with plant resistance factors should be used with caution
 - Eta could be calculated more simply and more accurately by using diurnal ground water fluctuations
-
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***Thank You for
Your attention***

The „oasis-effect“



Method II - Hydrological measurement program

The willow site

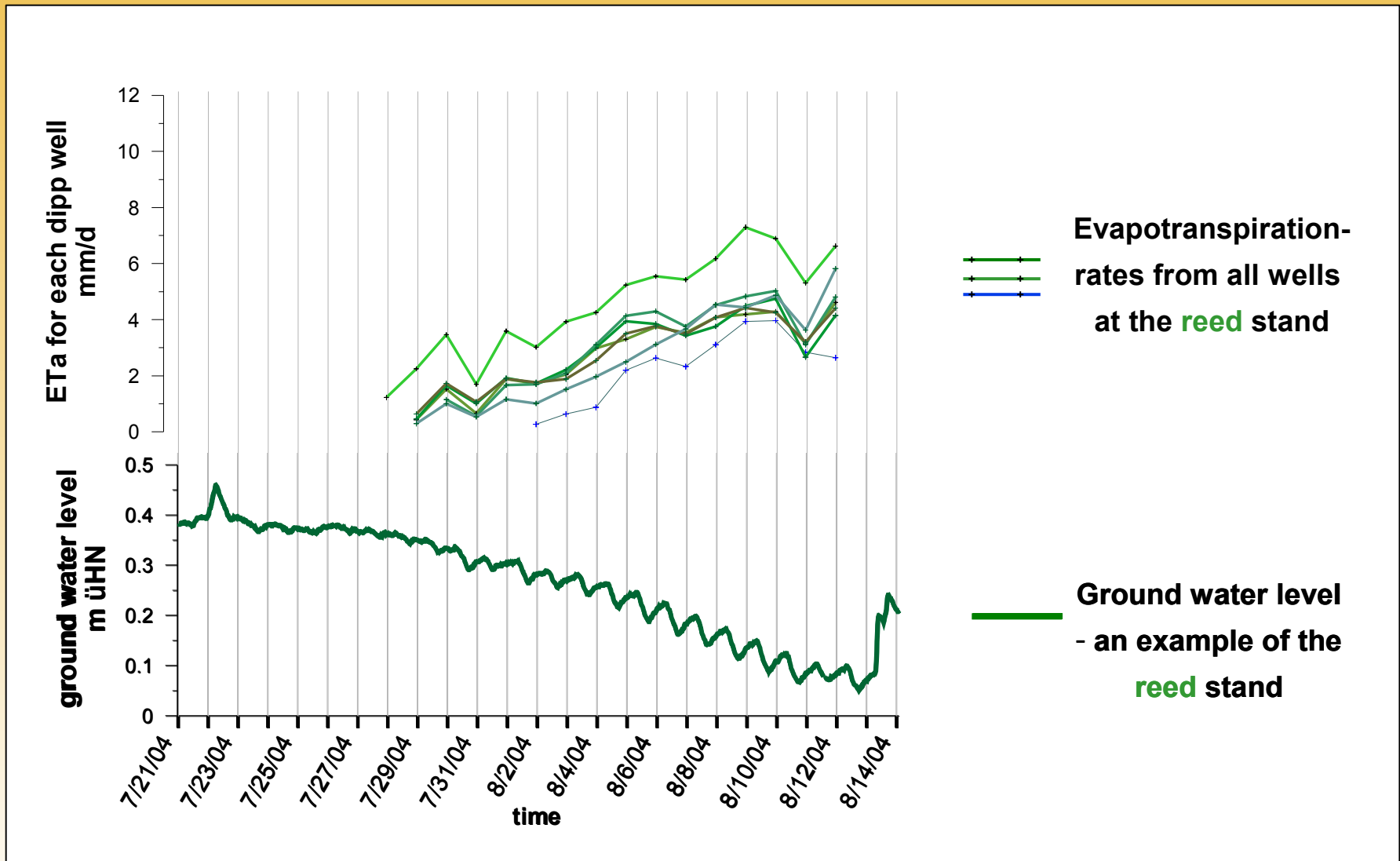


View to the willow stand



Groundwater dip well

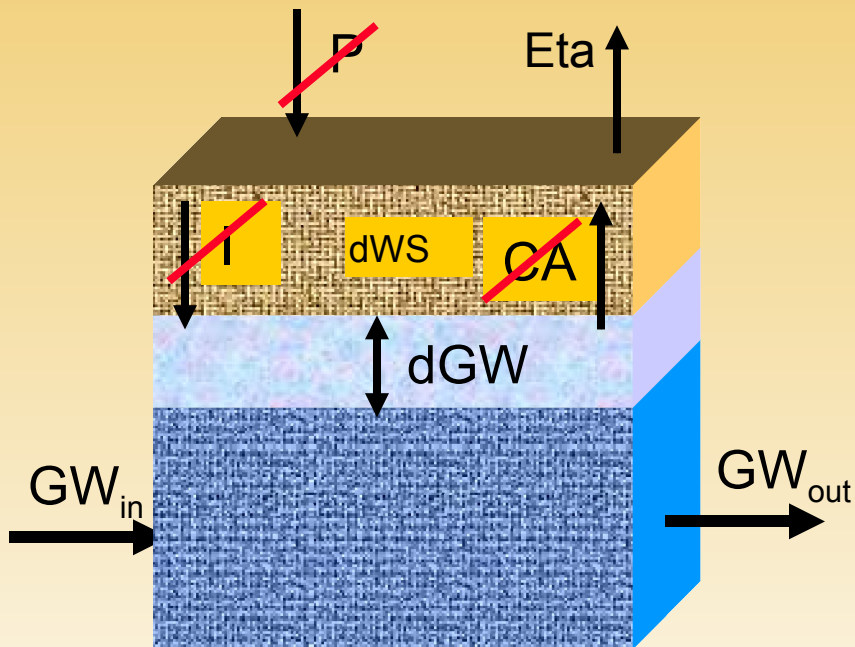
Results – The Draw down recharge method – at the reed stand



Method II – Hydrological measurement program

accounting equation:

$$E_a = \cancel{P} + \cancel{I} - \cancel{CA} - dWS - dGW + (GW_{in} - GW_{out})$$



P - precipitation

I - influent seepage

CA - capillary action

dWS - Water storage

dGW - ground water fluctuation

GW_{in} - inflow GW

GW_{out} - outflow GW

↓

$E_a = dGW$